

# A comparison of selected natural radionuclide concentrations in the thermal groundwater of Mszczonów and Cieplice with deep well water from Łódź city, Poland

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**Abstract.** A simple procedure using  $\alpha$  and  $\gamma$  spectrometry for the determination of natural radionuclides in two Polish geothermal water samples from Mszczonów and Cieplice, as well as in deep well groundwater from the city of Łódź, was described. The  $^{222}\text{Rn}$  radionuclide was directly extracted from  $10\text{ cm}^3$  water samples to a scintillator solution placed in scintillation vials and determined by a liquid scintillation (LSC) technique with  $\alpha/\beta$  separation. The activity concentrations of three radium nuclides:  $^{228}\text{Ra}$  ( $^{228}\text{Ac}$ ),  $^{226}\text{Ra}$  ( $^{222}\text{Rn}$ ),  $^{224}\text{Ra}$ , as well as  $^{210}\text{Pb}$  and  $^{228}\text{Th}$  were measured by  $\gamma$  spectrometry with an HPGe detector after their preconcentration on hydrated  $\text{MnO}_2$ .  $^{210}\text{Po}$  was deposited on a silver disc from dissolved  $\text{MnO}_2$  precipitate after  $\gamma$ -spectrometry analysis and measured by  $\alpha$  spectrometry. The concentration of  $^{210}\text{Pb}$  in the examined samples was below the detection limit of the method ( $< 6.2\text{ mBq/dm}^3$ ), whereas activity of its decay product  $^{210}\text{Po}$  was in the range  $0.35\text{--}1.4\text{ mBq/dm}^3$ . Higher activities of  $^{228}\text{Ra}$  ( $46.7\text{ mBq/dm}^3$ ),  $^{226}\text{Ra}$  ( $67.8\text{ mBq/dm}^3$ ) and  $^{224}\text{Ra}$  ( $22.4\text{ mBq/dm}^3$ ) were observed for the deep well water in Łódź in comparison to those for geothermal water samples from Cieplice and Mszczonów. However, concentrations of all the measured radionuclides were below World Health Organization (WHO) reference activities ( $0.1\text{ Bq/dm}^3$  for  $^{228}\text{Ra}$  and  $1\text{ Bq/dm}^3$  for  $^{226}\text{Ra}$  and  $^{226}\text{Ra}$ ).

**Key words:** thermal groundwater • natural radionuclides • ICP-MS • co-precipitation

## Introduction

It is expected that due to availability, and economic and environmental issues geothermal energy will play an important role in future worldwide energy budgets. However, because of the relatively high cost of deep drilling, the majority of the medium and low temperature geothermal water fields were in the last decade exploited for bathing, swimming pool heating and, to a lesser extent, for space heating [18]. On the other hand, an observed drinking water shortage or the necessity of arid land farming in many countries has resulted in increasing interest in using geothermal produced fluids for these purposes [13].

Now, in Poland there are several existing geothermal plants, and a few new geothermal installations are planned in the so-called Szczecin-Łódź geothermal district (Fig. 1).

Depending on factors of geothermal waters like their energetic potential (temperature of outcoming water) mineralization and chemical composition, they have been used almost exclusively in Poland for heating purposes, recreation and balneology [8, 14]. There are only three sites in Poland with sufficiently low mineralization thermal groundwater which can be, and in fact are, used as a source of drinking water. These places are Mszczonów, Cieplice and Słomniki. Thermal groundwaters from Mszczonów and Cieplice have a relatively low mineralization ( $\sim 0.5\text{ g/dm}^3$  and below  $1\text{ g/dm}^3$ , respectively) in comparison with waters

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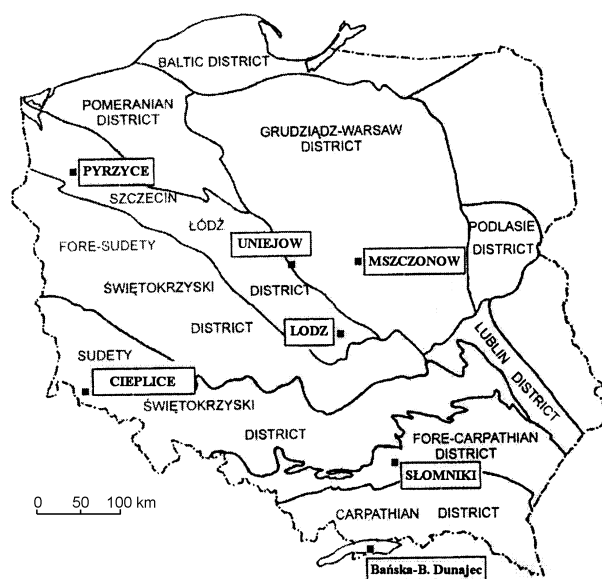


Fig. 1. Localization of the geothermal systems [14].

from other sites in Poland: 3 g/dm<sup>3</sup> in Bańska Biała or 5 g/dm<sup>3</sup> in Uniejow [15, 16]. The thermal water from the Mszczonów wellbore, with its temperature of 42.2°C, comes from 1700 m below ground level, with an average flow rate of 50 m<sup>3</sup>/h. In Słomniki water is drawn from 300 m below the surface (flow rate 50–100 m<sup>3</sup>/h), thus its temperature is only about 20°C [15, 19].

In the case of utilization of geothermal water for balneological or drinking water purposes, exact chemical studies including the determination of trace elements and radionuclide concentrations must be performed. Thermal groundwater used as drinking water additionally must fulfill very restrictive regulations of radionuclide content. According to the Polish state regulations [25], which are in accordance with EU Council Directive 98/83/EC [12] and WHO recommendations [32], the committed effective dose for members of the public from 1 year's consumption of drinking water (~ 730 dm<sup>3</sup>) should not exceed 0.1 mSv [32]. Consequently, the WHO guidance level values for activity concentrations in drinking water were calculated taking into account their dose conversion factors (mSv/Bq·dm<sup>-3</sup>) and corresponding activity concentrations, as shown in Table 1. The WHO protocol for radioactivity measurements in drinking water recommends, as the first step, determination of gross  $\alpha$  and gross  $\beta$  activity [30, 32]. When gross  $\alpha$  or gross  $\beta$  activity exceeds 0.5 and 1 Bq/dm<sup>3</sup>, respectively, determination of individual radionuclides must be performed. Water must not be used as drinking water when the measured activity

Table 1. WHO guidance levels of natural radionuclides in drinking water [32]

Radionuclide	Activity (Bq/dm <sup>3</sup> )
<sup>238</sup> U	10
<sup>226</sup> Ra	1
<sup>222</sup> Rn	100
<sup>210</sup> Pb	0.1
<sup>210</sup> Po	0.1
<sup>228</sup> Ra	0.1
<sup>228</sup> Th	1
<sup>224</sup> Ra	1

of an individual radionuclide is equal or higher than the guidance level values.

As is evident from Table 1, the most dangerous natural radionuclides in drinking water are <sup>210</sup>Po, <sup>210</sup>Pb and <sup>228</sup>Ra, whose permissible activities in water must not exceed 0.1 Bq/dm<sup>3</sup>. These three nuclides belong to a natural decay series and occur in very low concentrations in surface water. However, the activity of these and other natural radionuclides in underground waters depends on the activity of soil and rocks which surround groundwater resources. In most cases of groundwater samples <sup>238</sup>U concentrations are lower than its progeny radionuclides, what is caused by low transfer of this nuclide to water from bedrock [31].

Concentrations of radionuclides in natural water are usually rather low [9] and before their determination a preconcentration step must be applied. In the last decades, for the determination of low activities of uranium, thorium and other long-lived radionuclides, the inductively coupled plasma with mass spectrometry (ICP-MS) method was successfully applied [1–5, 20, 29]. However, because of the high equipment costs for this method, radiometric procedures are still widely used for radionuclide determination in water. An effective and simple method of radionuclide concentration from water samples is their co-precipitation with hydrated manganese dioxide [28]. This method was successfully applied for uranium, thorium and radium radionuclides as well as for <sup>210</sup>Po and <sup>210</sup>Pb. After preconcentration of radionuclides, they can be measured by  $\gamma$  or  $\alpha$  spectrometry methods or also by LSC methods [11, 21, 22, 26, 27].

Although concentrations of radon, uranium and radium isotopes in Polish groundwaters and in bottled mineral waters have been well documented [6, 17, 22–24, 27], until now the activities of some radionuclides were reported for geothermal water from the Uniejow plant only [7].

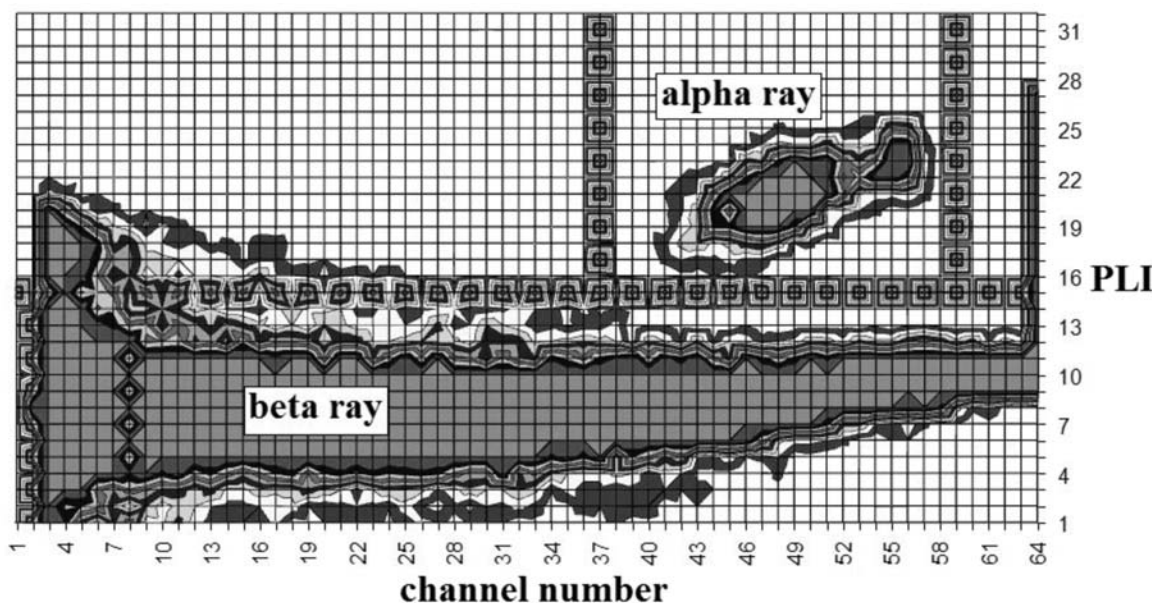
The aim of these studies was the elaboration of a simple radiochemical procedure for determination of the activity of principal natural radionuclides occurring in thermal groundwater (Mszczonów, Cieplice). For comparison, the activities of the same radionuclides were determined in one of the several deep wells (Lower Cretaceous geological formation) connected to the Łódź city municipal water system with mineralization ~ 0.2 g/dm<sup>3</sup>. Additionally, natural radionuclides in Łódź tap water, which is a mixture of the deep well groundwater with surface water from Sulejów reservoir, were also measured.

## Experimental

### Materials and methods

#### Sampling

Thermal groundwater samples from Mszczonów were collected from Geotermia Mazowiecka S.A. and in the case of Cieplice directly from the public well situated in the centre of the city. Drinking water samples from Łódź were taken from two locations: deep well water from the Dąbrowa borehole and from the Łódź municipal water network in the centre of the city. All chemicals



**Fig. 2.** Energy spectrum of  $^{222}\text{Rn}$  as pulse shape discrimination (PSD) plot (2D) x-axes  $\times 16 =$  channel number, y-axes  $\times 32 =$  PLI (pulse length index).

were of analytical grade purity, used as provided. Standard reference material IAEA-327 was supplied by the International Atomic Energy Agency (IAEA).

#### Sample preparation

For the determination of  $\gamma$ -emitting radionuclides, a preconcentration step on hydrated manganese dioxide from  $5 \text{ dm}^3$  of water samples was applied, according to procedure described elsewhere [7].

#### Radioactivity measurements

$^{222}\text{Rn}$  activity was measured by means of a liquid scintillation technique, after its direct extraction in scintillation vials from  $10 \text{ cm}^3$  of analysed water using  $10 \text{ cm}^3$  of a Ultima Gold F scintillation cocktail. Samples counting were performed for 40,000 s using a BetaScout (PerkinElmer) liquid scintillation counter with  $\alpha/\beta$  separation option, after at least 3 h necessary for establishing radioactive equilibrium of  $^{222}\text{Rn}$  and its progeny ( $^{218}\text{Po}$  and  $^{214}\text{Po}$ ).

Optimal  $\alpha/\beta$  separation conditions in order to eliminate the overlapping of the  $\beta$  pulses to the  $\alpha$  region were chosen for this device after preliminary counting and setting the discrimination levels. Figure 2 shows that complete distinction of  $\alpha$  from  $\beta$  pulses is possible for the Ultima Gold F scintillation solution for pulse length index (PLI) equal to 15. The observed background in the chosen  $\alpha$  region was very low  $< 0.2 \text{ cpm}$ .

The activity of  $\gamma$ -emitting radionuclides adsorbed on a  $\text{MnO}_2$  precipitate was determined using a Canberra spectrometry system with an HPGe detector with a relative efficiency of 25%, according to the procedure described elsewhere [7].

Because of the relatively high detection limit for  $^{226}\text{Ra}$  determination from its  $\gamma$  line of 185.6 keV ( $8.5 \text{ mBq/dm}^3$  according to the Curie criterion [10]) for the Mszczonów thermal groundwater  $^{226}\text{Ra}$  was additionally determined by the liquid scintillation method after extraction of the  $^{222}\text{Rn}$  from  $0.5 \text{ dm}^3$  samples [6].

The samples of manganese dioxide precipitate after  $\gamma$ -spectroscopy measurements were digested with 3% (w/w) solution of hydrogen peroxide in  $70 \text{ cm}^3$  of 1 M HCl. The obtained solutions were transferred to a teflon vial, and spontaneous deposition of  $^{210}\text{Po}$  on a silver disc was performed for about 8 h at a temperature of about  $80^\circ\text{C}$ .  $^{210}\text{Po}$  after its deposition on the silver plate was counted using the  $\alpha$ -spectrometry system with a passivated implanted planar silicon (PIPS) semiconductor detector (Canberra 7401VR alpha spectrometer) for 2–3 days.

#### Quality assurance

The accuracy of the analytical procedure was determined in an independent experiment by checking the recovery of  $^{210}\text{Po}$ ,  $^{210}\text{Pb}$ ,  $^{226}\text{Ra}$ ,  $^{228}\text{Ra}$  and  $^{238}\text{U}$  radionuclides from 1 g of the mineralized standard reference material: SRM-IAEA 327 "Radionuclides in soil". The results are presented in Table 2.

**Table 2.** Chemical recoveries of the determined radionuclides

Nuclide	Reference value (mBq/g)	Measured value (mBq/g)	Recovery (%)
$^{210}\text{Po}$	49.84	$49.33 \pm 3.95$	99.0
$^{210}\text{Pb}$	49.84	$48.08 \pm 4.49$	96.5
$^{226}\text{Ra}$	33.89	$34.04 \pm 11.51$	100.5
$^{228}\text{Ra} (^{228}\text{Ac})$	38.7	$35.21 \pm 3.65$	91.0
$^{238}\text{U} (^{234}\text{Th})$	32.8	$32.95 \pm 5.07$	100.5

**Table 3.** Concentration of the main natural radionuclides in water samples

Nuclides	$\gamma$ line (keV)	$A_{\text{Mszczonow}}$ (mBq/dm <sup>3</sup> )	$A_{\text{Cieplice}}$ (mBq/dm <sup>3</sup> )	$A_{\text{Lodz tap}}$ (mBq/dm <sup>3</sup> )	$A_{\text{Lodz deep well}}$ (mBq/dm <sup>3</sup> )
<sup>238</sup> U( <sup>234</sup> Th)	63.3	< 4.7	≤ 4.7	< 4.7	< 4.7
<sup>226</sup> Ra	185.6	14.8 ± 4.3	50.0 ± 5.3	< 8.5	67.8 ± 6.5
<sup>226</sup> Ra( <sup>222</sup> Rn)	LSC	18.0 ± 0.6	–	–	–
<sup>222</sup> Rn	LSC	1930 ± 34	9100 ± 560	1093 ± 20	2860 ± 190
<sup>210</sup> Pb	46.5	≤ 6.2	< 6.2	< 6.2	< 6.2
<sup>210</sup> Po	$\alpha$ counting	1.4 ± 0.1	0.4 ± 0.1	< 0.2	≤ 0.2
<sup>228</sup> Ra( <sup>228</sup> Ac)	911.2	15.3 ± 1.4	10.4 ± 1.5	< 3.3	46.7 ± 2.8
<sup>228</sup> Th	84.3	29.1 ± 5.9	37.9 ± 6.0	< 12.6	116.4 ± 9.4
<sup>224</sup> Ra	241.0	≤ 5.7	< 5.7	< 5.7	22.7 ± 4.5

## Results

The results of activity concentration of measured radionuclides in different water samples are presented in Table 3.

The measured activities, except for <sup>222</sup>Rn, were generally very low. It is interesting that the activity of <sup>224</sup>Ra, <sup>226</sup>Ra and <sup>228</sup>Th in the deep well water samples coming from the Lower Cretaceous geological formation were higher than those for thermal groundwater from Cieplice and Mszczonów, but remarkably lower than the corresponding activities for the previously determined thermal groundwater for Uniejów [7]. The slightly higher temperature of the Mszczonów (42°C) and Cieplice (30°C) thermal groundwater in comparison with deep well water (4°C) does not influence the mobility of these radionuclides from the adjacent geological formations, and the low concentrations of natural radionuclides in these waters are caused by a remarkable input of infiltrated surface water to their underground reservoirs.

The drinking tap water in the Łódź municipal network is a mixture of water coming from several deep wells and surface water pumped from the Sulejów water reservoir. Therefore, concentrations of natural radionuclides in this kind of water (additionally purified in the water supply plants) is extremely low (except for <sup>222</sup>Rn) below the detection limits.

Among the radionuclides in all the investigated water samples, the highest activities were observed for <sup>222</sup>Rn radionuclide. Depending on the origin of water, the <sup>222</sup>Rn activity varied between 1093 mBq/dm<sup>3</sup> and 9100 mBq/dm<sup>3</sup> for the Cieplice samples (Table 3). The last values are in close accord with the data for radon concentration in the groundwaters of the Sudety Mountains where elevated activities of this radionuclide were observed [24].

It is worth underlining that the concentrations of all measured radionuclides were below the recommended WHO values.

## Conclusion

1. Co-precipitation of the natural radionuclides with hydrated manganese dioxide from a large volume of water samples can be used for the sample preparation before  $\gamma$ -spectrometry counting.
2. Activities of natural radionuclides from <sup>238</sup>U and <sup>232</sup>Th decay series in thermal groundwater from Mszczonów and Cieplice, and deep well water in

Łódź are below the threshold limit values according to EU and WHO directives for drinking water.

3. The low levels of natural radionuclides in the thermal groundwater from Cieplice and Mszczonów, in contrast to the Uniejów region thermal groundwater, indicate a remarkable input of infiltration of surface waters with a low mineralization to their underground reservoirs.
4. From the radiological points of view both thermal groundwaters from Mszczonów and the Cieplice public well can be used as drinking water.

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